

A stress destruction stage field model on behalf of hydrofracturing of crawling ice shelves and glaciers

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ABSTRACT

There is a need for computational models capable of predicting meltwater-assisted crevasse growth in glacial ice. Mass loss from glaciers and ice sheets is the largest contributor to sea-level rise and iceberg calving due to hydrofracture is one of the most prominent yet less understood glacial mass loss processes. To overcome the limitations of empirical and analytical approaches, we here propose a new phase field-based computational framework to simulate crevasse growth in both grounded ice sheets and floating ice shelves. The model incorporates the three elements needed to mechanistically simulate hydrofracture of surface and basal crevasses: (i) a constitutive description incorporating the non-linear viscous rheology of ice, (ii) a phase field formulation capable of capturing cracking phenomena of arbitrary complexity, such as 3D crevasse interaction, and (iii) a poro-damage representation to account for the role of meltwater pressure on crevasse growth. A stress-based phase field model is adopted to reduce the length-scale sensitivity, as needed to tackle the large scales of iceberg calving, and to adequately predict crevasse growth in tensile stress regions of incompressible solids. The potential of the computational framework presented is demonstrated by addressing a number of 2D and 3D case studies, involving single and multiple crevasses, and considering both grounded and floating conditions. The model results show a good agreement with analytical approaches when particularised to the idealised scenarios where these are relevant. More importantly, we demonstrate how the model can be used to provide the first computational predictions of crevasse interactions in floating ice shelves and 3D ice sheets, shedding new light into these phenomena. Also, the creep-assisted nucleation and growth of crevasses is simulated in a realistic geometry, corresponding to the Helheim glacier. The computational framework presented opens new horizons in the modelling of iceberg calving and, due to its ability to incorporate incompressible behaviour, can be readily incorporated into numerical ice sheet models for projecting sea-level rise.

KEYWORDS: Fracture Mechanics, Structural Integrity, Hydrogen Embrittlement, Nondestructive Evaluation

1.0 INTRODUCTION

Ice sheets are large masses of glacial ice that inundate the surrounding landscape in Greenland and Antarctica today, and many other regions during ice ages [1-23]. These act as enormous stores of freshwater – containing approximately 70% of the planet's supply [24-38] – that assist in regulating a stable global climate, through maintaining global ocean-water levels and controlling surface temperatures by reflecting solar radiation due to its high albedo properties [39-47]. Ice sheets thin toward their margins, and if these are located in marine settings, they will form floating extensions known as ice shelves, which act to provide resistive buttressing to downslope flow and reduce the flux of grounded ice to the ocean. However, increasing global temperatures as a result of carbon emissions has led to higher rates of ablation than accumulation, resulting in ice shelf and ice sheet thinning in some key areas where ice-sheet instability may follow [48-65]. Surface and basal crevasses can form within ice sheets as a consequence of ongoing deformations within the ice. These are deep crack-like defects that can propagate in an unstable manner and lead to large-scale iceberg calving events, and in extreme cases the catastrophic break up of ice shelves. The frequency of these events has grown in recent decades, beginning with the disintegration of Larsen A (1995) and Larsen B (2002) ice shelves, and more recently significant surface melting and iceberg calving on Larsen C (2017), Pine Island and Thwaites (2018–2020), and Conger (2022) ice shelves [66-79]. Fracture within ice shelves can result in a loss of resistance to down slope glacial flow, leading to ice-sheet thinning, additional flotation of grounded ice and, thus, potentially irreversible grounding line retreats [1-19]. Deposition of grounded glacial ice into the ocean is one of the leading contributors to sea level rise [20-34], having direct implications within this Century on low-lying coastal regions through flooding, increased extreme environmental events, degradation of farmland and loss of habitat, among others. A key driving factor

for their stability is the production of surface meltwater as a result of elevated surface temperatures [35-49]. When ice shelves and glaciers melt, meltwater flows down-slope into surface crevasses, causing additional tensile stresses to form within the crevasse. This leads to crevasse instability, and with sufficient meltwater, the crevasse can propagate through the full thickness of the ice column. This process is generally referred to in the glaciological literature as hydrofracture [50-67]. A recent study by Lai et al. found that approximately $60 \pm 10\%$ of Antarctic ice shelves provide significant buttressing to downslope flow and are vulnerable to meltwater driven hydrofracture, highlighting the significance of studying the formation and propagation of crevasses in glaciers [68-79]. An illustration of a grounded ice sheet, transitioning to a crevassed floating ice shelf is shown in Fig. 1.

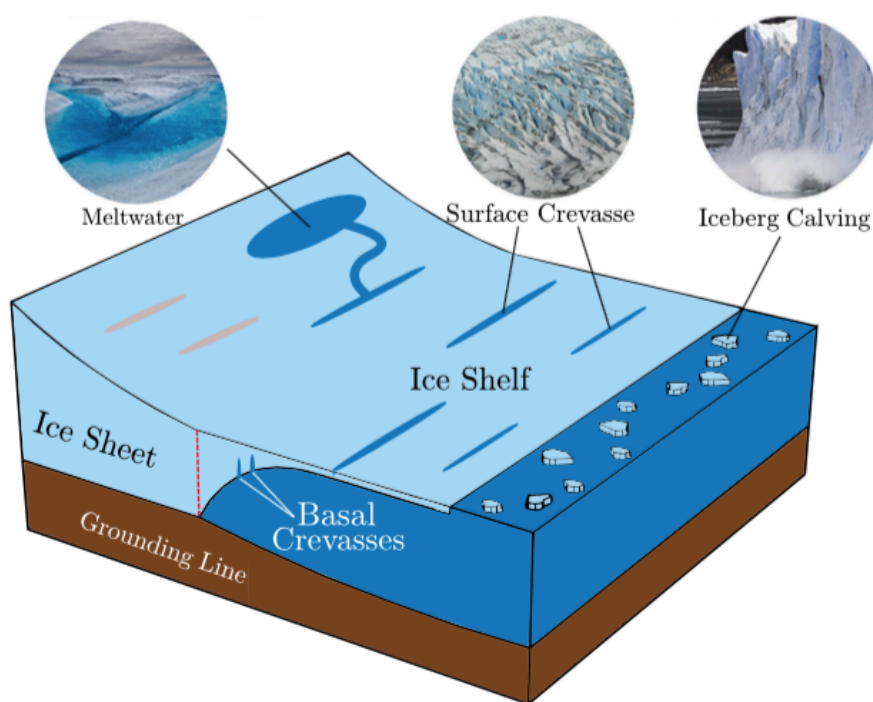


Fig. 1. Illustration of a grounded ice sheet and a floating ice shelf, containing both surface and basal crevasses, and with calving events occurring at the terminus.

integrating over the crevasse depth for the normal tensile stress, the lithostatic compressive stress, and the meltwater pressure. In order for crevasses to stabilise, the net stress intensity factor K_{net} must be equal to the material's fracture toughness K_c . However, these analytical approaches have well-known limitations, such as (i) idealised scenarios and boundary conditions are assumed; (ii) creep effects, resulting from the continual movement of glaciers under their own weight, are neglected; and (iii) crevasse interaction cannot be captured. Recently, computational methods have been used to predict crevasse growth and iceberg calving events. Local and non-local continuum damage mechanics formulations have been presented to predict ice sheet fracture [1-22]. These works have overcome some of the limitations intrinsic to analytical approaches, but often at the cost of using empirical parameters. Variational phase field fracture models offer an alternative approach, enabling the simulation of realistic conditions (3D geometries, multiple interacting crevasses, etc.) and providing a connection to fracture mechanics theory. Phase field fracture models have gained remarkable popularity in recent years due to their ability to predict complex cracking phenomena including crack bifurcation, coalescence and nucleation from arbitrary sites [23-35]. New phase field-based formulations have been presented for dynamic fracture [36-47], ductile damage [48-59], environmentally assisted cracking [30,31], fatigue crack growth [17-33], hydraulic fracture [60-67], and battery degradation [4-37]; among other (see Refs. [68-79] for an overview). In this work, we aim at extending the success of phase field fracture models to the area of glacier crevassing and iceberg calving. To this end, a new phase field formulation is presented capable of capturing the creep behaviour of glacial ice and the role of fluid pressure in driving crevasse growth. Also, for the first time, crevasse interaction is predicted in both 2D and 3D. Very recently, Sun et al. [40-56] used a phase field approach to predict hydrofracture in 2D linear elastic glaciers, assuming compressible

behaviour and disregarding creep effects. Unlike them, we base our framework on a stress-based phase field fracture formulation, which offers several advantages in the context of hydrofracturing of glacier crevasses. First, strain energy-based approaches are unsuited for incompressible rheologies. This is not only important due to the incompressible nature of glacial ice, but also because it hinders its integration into large-scale computational models for ice sheet evolution and sea level rise, which assume incompressible flow (see, e.g., the Community Ice Sheet Model (CISM) [57-68]). Second, ice-sheet fracture is driven by tensile stresses and not strains, with crevasses propagating solely in regions where the net longitudinal stress is positive [42]. This is naturally accounted for in a stress-based phase field model, while requiring a particular ad hoc split in strain energy-based formulations [69-79]. Third, a phase field length-scale insensitive driving force can be defined, enabling the use of coarser meshes, a key enabler given the large scales involved. These advantages provide further motivation for this work, presenting the first stress-based phase field computational framework for hydrofracturing of creeping glaciers and ice shelves [1-18]. The rest of the paper is outlined as follows. The theoretical and computational framework presented is described in Section 2. The model is then used in Section 3 to predict hydrofracturing in case studies of particular interest. First, the propagation of single crevasses in grounded ice considering both linear and non-linear rheologies is investigated. A parametric study is conducted to assess the role of relevant material parameters, seawater level and meltwater depth. Second, we simulate the growth of a field of densely spaced crevasses in a grounded glacier, comparing against the predictions of Nye's zero stress model [19-36]. Third, the growth of basal and surface crevasses (and their interaction) is for the first time simulated for a floating ice shelf, using appropriate Robin boundary conditions. Fourth, the combined creep-phase field fracture model is used to predict the nucleation and growth of crevasses in a realistic geometry, corresponding to the Helheim glacier. Finally, we provide the first 3D analysis of crevasse propagation in ice sheets. Concluding remarks end the manuscript in next section [37-54].

2.0 METHODOLOGY

In this section, we present our computational framework, which encompasses the three elements that are needed to resolve the hydrofracture process taking place in ice sheets; namely, the viscoplastic behaviour of ice, the propagation of meltwater-filled crevasses, and the role of meltwater pressure on crevasse propagation. These are modelled by means of Glen's flow law [24-41], a stress-based phase field description of fracture [1-18], and a meltwater-ice poro-damage model [42-58], respectively. Fig. 2 illustrates upon a single crevasse the mechanistic and modelling assumptions of our framework. In the following, we present the kinematics of the problem (Section 2.1), formulate the energy functionals, particularise the model upon suitable constitutive choices (Section 2.3), and briefly describe the finite element implementation (Section 2.4). Throughout, the formulation refers to a body occupying an arbitrary domain $\Omega \subset \mathbb{R}^n$ ($n \in [1, 2, 3]$), with an external boundary $\partial\Omega \subset \mathbb{R}^{n-1}$ with outwards unit normal \mathbf{n} .

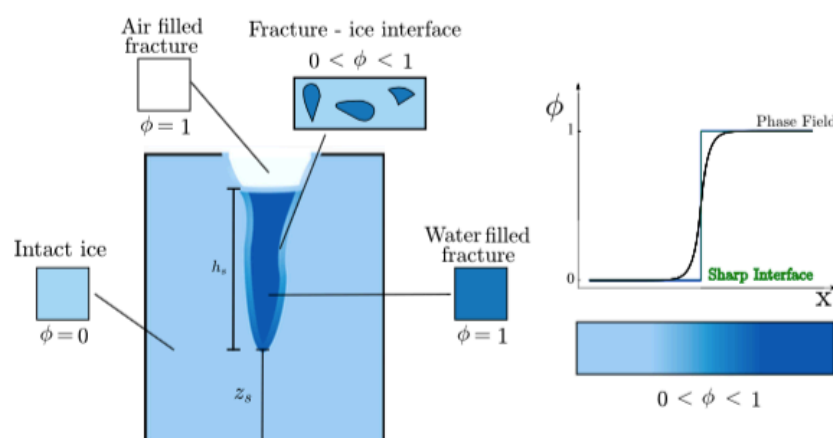


Fig. 2. Schematic diagram of a meltwater filled crevasse in glacial ice, illustrating the intact phase ($\phi = 0$), fully cracked phase ($\phi = 1$) and transition phase ($0 < \phi < 1$). In the damaged and transition phases, there is a hydrostatic pressure contribution to damage arising from the meltwater. Relevant to the poro-damage part of the model, h_s denotes the meltwater depth, and z_s is the distance between the glacier base and the bottom of the crevasse, with z being the vertical height.

3.0 RESULT

In this section, we present a series of 2D and 3D numerical examples, aimed at capturing the propagation of surface and basal crevasses within grounded glaciers and floating ice shelves. For 2D examples, we consider an idealised rectangular glacier of length $L = 500$ m and height $H = 125$ m, under the assumption of plane strain conditions. For simplicity, we neglect lateral shear and restrict the domain to a flow line near the terminus with x and z representing the along-flow and vertical coordinates [33-47]. Gravitational load due to self-weight is applied as a uniform body force in the z -direction with a magnitude of $-\rho_i g$. We also consider the surface meltwater pressure p_w within a crevasse using the poro-damage approach presented in Eq. (19). A Neumann-type traction is applied normal to the ice-ocean interface at the terminus, with the hydrostatic ocean-water pressure varying linearly with depth and a magnitude of $-\rho_s g (h_w - z)$. Boundary conditions that are specific to the grounded glacier and floating ice shelf cases are discussed in Sections 3.1 and 3.3, respectively. Our simulations deal with glacial ice, whose material properties are given in our literature review, along with the densities of seawater and meltwater. The strength σ_c magnitude is chosen to be an intermediate magnitude within the experimentally reported values of the critical fracture stress in glacial ice, which are in the range 0.08–0.14 MPa [48-71]. An estimate of the phase field length scale, which plays a negligible role in this model, can be obtained through the Hillerborg et al. relation, which for plane strain reads: $\ell = (1 - \nu^2)Kc^2/\sigma_c^2$. Considering the toughness of glacial ice ($Kc = 0.1$ MPa $\sqrt{\text{m}}$), this gives a magnitude of $\ell = 0.625$ m, which is the value adopted here (unless otherwise stated) [58-79]. To attain mesh-independent results, the characteristic element size along the crevasse propagation region is always chosen to be at least 5 times smaller than the phase field length scale ℓ .

Table 1

Material properties assumed in this work (unless otherwise stated). The values are chosen to characterise the behaviour of glacial ice, with the subscript number denoting the relevant reference.

Material parameter	Magnitude
Young's modulus, E [MPa]	
Poisson's ratio, ν [–]	
Density of glacial ice, ρ_i [kg/m ³]	
Density of meltwater, ρ_w [kg/m ³]	
Density of seawater, ρ_s [kg/m ³]	
Fracture toughness, K_c [MPa $\sqrt{\text{m}}$]	
Critical fracture stress, σ_c [MPa]	
Creep exponent, n [–]	
Creep coefficient A [MPa ^{$-n$} s ^{-1}]	

4.0 CONCLUSIONS

We studied the role of hydrogen in strength degradation of micro-architected materials. We began by analyzing the Reynold's transport equation for hydrogen diffusion problem and obtained the governing equations for hydrogen concentration under quasi-static loading. We then used the concept of elastoplastic homogenization to bridge the scales and derive the material effective properties. We implemented these theories in a numerical scheme for decoupled diffusion-deformation analysis, in which we update the homogenization algorithm with hydrogen concentration and flow stress through UHARD subroutine and implicit formulae. We applied the model to cubic (with $\rho_r = 10\%$, 20% and 30%) and BCC (with $\rho_r = 20\%$) unit-cells along with PBC and characterized their macroscopic hydrogen degradation laws. Each micro-architected material has its unique failure loci depending on cell architecture, but one can provide a single expression to describe the behavior of a cubic unit-cell over a range of relative densities. Also, it turned out that the role of trap hydrogen in embrittlement of micro-architecture materials is negligible, especially when the base material has low ductility. For load bearing applications and in presence of hydrogen, the cubic material outperforms the BCC because: (1) at equal hydrogen contents, its maximum strength is higher than BCC, and (2) it undergoes a less severe hydrogen degradation as compared to the BCC. On the other hand, the BCC material has higher ductility which is desirable for certain applications. Generally, micro-architected materials made of high-strength steels, e.g. AISI 4135, are prone to brittle fracture in the presence of hydrogen. This agrees with what was observed in bulk specimens, e.g. The developed computational scheme is generic and applicable to any periodic micro-architected material and is an efficient tool for assessment of hydrogen embrittlement. We proposed the design methodology and numerically characterized the macroscopic hydrogen degradation laws for cubic and BCC unit-cells. The experimental validation of the results on hydrogen embrittlement in architected materials is suggested as a part of future studies.

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